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The DOI for this manuscript is

DOI:10.2151/jmsj.2024-005

J-STAGE Advance published date: November 2nd, 2023

The final manuscript after publication will replace the preliminary version at the above DOI once it is available.

# Roles of August Kuroshio SST Anomaly in Precipitation Variation during September over Central China

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Sep 27, 2023

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**Abstract**

29       Sea surface temperature anomaly (SSTA) in the Kuroshio region near the East China  
30 Sea (K-ECS) during August has been found to be closely related to the September  
31 precipitation anomaly over Central China (CC) in this study. The significant causality is  
32 identified from SSTA in K-ECS to anomalous rainfall over CC. When a negative SSTA occurs  
33 in K-ECS during August, its cooling effect on the overlying atmosphere lasting from August  
34 to September promotes the formation of anomalous anticyclonic circulation below 600 hPa  
35 above K-ECS and the southeast of China by weakening the local convection. The southerly  
36 in the west of this anomalous circulation transports more water vapor into CC. Eventually, the  
37 convergence of moisture provides favorable conditions for the generation of precipitation here.  
38 The responsive characteristics of the anomalous atmospheric circulation during September  
39 to a negative SSTA in K-ECS in August can be confirmed through numerical experiments.  
40 The above important long-term relationship suggests that SSTA in K-ECS during August  
41 could serve as a valuable predictor for September precipitation over CC.

42 **Keywords** precipitation; Central China; sea surface temperature; Kuroshio near the East  
43 China Sea

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## 49 **1. Introduction**

50       Precipitation is one of the most important meteorological variables in the study of climate  
51 change. It can induce severe meteorological disasters and further impact human life,  
52 socioeconomy and ecological environment (Niu and Li, 2008; Barriopedro et al., 2012; Yang  
53 et al., 2012; Zhang et al., 2012; Zhang et al., 2014; Zhou et al., 2018; Zhang et al., 2019).  
54 Autumn (SON: from September to November) is the key period for crop maturity, harvest and  
55 sowing winter wheat in China. The severe rainfall and drought in this season frequently cause  
56 enormous damage to agricultural production (Barriopedro et al., 2012; Zhang et al., 2012;  
57 Zhang et al., 2014; Zhou et al., 2018). Therefore, considerable attentions have been drawn  
58 to the causes of autumn precipitation.

59       El Niño–Southern Oscillation (ENSO), the strongest signal of interannual variations in  
60 the air-sea coupled system, has been extensively revealed by previous studies to induce  
61 autumn precipitation anomalies in China. The eastern-Pacific and central-Pacific types of El  
62 Niño change the transport path of tropospheric moisture and the location of anomalous water  
63 vapor divergence by affecting the Hadley circulation and the Western North Pacific  
64 Subtropical High (WNPSH), which further result in opposite precipitation anomalies in the  
65 southern China (Zhang et al., 2013; Gu et al., 2014; Zhang et al., 2014; Hu et al., 2018; Yuan  
66 and Wang, 2019). Additionally, the effects of La Niña on the subtropical and even midlatitude  
67 atmospheric circulations can lead to anomalous precipitation in the south and north of China,  
68 respectively (Feng and Wang, 2018; Yuan and Wang, 2019). Moreover, as the high-degree  
69 synergic air-sea interactions, both ENSO and Indian Ocean Dipole can cooperatively impact

70 boreal autumn rainfall in the eastern China by altering atmospheric circulation over the  
71 Western North Pacific (WNP) and the East Asian jet stream (Xiao et al., 2015; Xu et al., 2016;  
72 Li and Zhao, 2019). For the specific extreme event, apart from the involvement of ENSO, the  
73 long-lasting Madden-Julian Oscillation also plays an important role in autumn precipitation in  
74 China (Qi et al., 2021).

75 In addition to the influence of tropical sea surface temperature (SST) anomalies (SSTAs),  
76 numerous studies have focused on the role of mid-high latitude signals. From the upstream  
77 of China, the sea ice anomaly in Barents-Kara Seas yielding an eastward Rossby wave train  
78 and the North Atlantic Oscillation stimulating a downstream Atlantic-Eurasian teleconnection  
79 can induce autumn rainfall in the central-eastern China (Xu et al., 2013; Shen et al., 2019).  
80 As an important circulation system over China, the East-Asian trough, modulated by the  
81 phase of Pacific Decadal Oscillation or Asian-Pacific Oscillation, is one of the main factors  
82 affecting autumn precipitation north of the Yangtze River (Qin et al., 2018; Wang et al., 2019;  
83 Lin et al., 2021). Besides, as a downstream signal, the generation of autumn tropical cyclones  
84 over WNP is accompanied by abundant moisture transport, which further enhances  
85 convective precipitation in the southern China (Wu et al., 2005; Hong et al., 2016; Hu et al.,  
86 2017).

87 Most studies on autumn precipitation in China have concerned the entire season, but  
88 there are some differences in precipitation in different months of autumn. The spatial  
89 distribution of rainfall in September is distinct from that in the remaining two months in autumn  
90 (Fig. 1). Comparatively, the standard deviation of accumulated precipitation over Central

91 China is significantly larger in September. Apart from precipitation, the atmospheric  
92 circulation regime in September is also different from that in the other two months, which is  
93 characterized by the abrupt southward migration of the East Asian jet and WNPSH after  
94 September (Yeh et al., 1958; Liu et al., 2022). The distinctiveness of September atmospheric  
95 variables in autumn suggests that the anomalous precipitation in this month is worth studying  
96 separately. Given the noticeable variability (Fig. 1e), the present study aims to explore the  
97 mechanism of the precipitation anomaly over Central China during September.

98 Kuroshio is an important western boundary current in the Pacific Ocean, characterized by  
99 a high ocean current velocity and a prominent ocean temperature front. Regarding  
100 anomalous precipitation in China, extensive researches have focused on the impacts of SST  
101 near the Kuroshio region. Meiyu in China (known as Baiu in Japan and Changma in Korea)  
102 is a typical episode of the East Asian rainy season (Liu et al., 2020). The variations of  
103 anomalous atmospheric circulation, influenced by the SST in the Kuroshio region near the  
104 East China Sea (ECS) through changing air temperature and humidity, can modulate the  
105 Meiyu rainfall (Matsumura and Horinouchi, 2023; Gan et al., 2019; Xu et al., 2018; Sasaki  
106 and Yamada, 2018; Minobe and Takebayashi, 2015; Sasaki et al., 2012). In warm season,  
107 the SSTA along the Kuroshio region (such as the region near the East China Sea and  
108 Kuroshio Extension) could lead to the geopotential height anomalies through altering lower  
109 atmospheric temperature, which affects the variations of WNPSH and thus results in the  
110 anomalous precipitation over the eastern China (Guan et al., 2019; Geng et al., 2022; Wang  
111 et al., 2023). Additionally, the Kuroshio thermal effect adjacent to the ECS can result in

112 anomalous planetary boundary layer height gradient through altering vertical mixing, which  
113 leads to plentiful vertical moisture transport and further enhances rainfall in the southern  
114 China during winter (Bai et al., 2023, 2020). In addition to the synchronous impact, several  
115 studies have indicated the importance of spring Kuroshio SST on summer precipitation in the  
116 eastern China, in which the teleconnection pattern acts as a bridge to establish the  
117 connection between the two variables (Fang et al., 2018; Ying et al., 2017).

118 Considering the essential influence of Kuroshio SST near the ECS on anomalous  
119 precipitation in China, it is natural to wonder whether the September rainfall anomaly over  
120 Central China can be affected by the Kuroshio SST? If so, what physical mechanisms are  
121 included in this process? To answer these questions, the remainder of this paper is organized  
122 as follows. In Section 2, data and methods are briefly introduced. The possible physical  
123 mechanisms of the precipitation over Central China in September are analyzed in Section 3.  
124 Finally, the summary and discussions are given in Section 4.

## 125 **2. Data and Methods**

126 In this study, the dataset of gridded monthly precipitation (Version 2.0), with a horizontal  
127 resolution of  $0.5^\circ \times 0.5^\circ$ , is kindly provided by the National Meteorological Information Center  
128 (NMIC) of the China Meteorological Administration (CMA). It is based on precipitation data  
129 from 2472 weather stations in China. The monthly SST at a  $0.25^\circ$  resolution, obtained from  
130 the Optimum Interpolation SST Version 2.1 (OISST; Reynolds et al., 2007; Huang et al.,  
131 2021), is employed for exploring the effect of anomalous oceanic signals. To analyze  
132 atmospheric dynamics, the monthly atmospheric variables, including geopotential height,

133 specific humidity, horizontal wind and vertical velocity, are taken from the fifth-generation  
 134 European Center for Medium-Range Weather Forecasts (ECMWF) reanalysis of the global  
 135 climate (ERA5; Hersbach et al., 2019) on a  $0.25^\circ$  grid. To ensure the consistency of analysis,  
 136 the time period of all variables has been selected from 1982 to 2021. Their anomalies are  
 137 formed by removing the long-term mean seasonal cycle and linear fit.

138 Given that the statistical relationship cannot reflect the causal relationship, a causality  
 139 analysis method based on Liang-Kleeman information flow theory (Liang 2014, 2015) is used  
 140 to unravel the causality between time series, which is applicable to both linear and nonlinear  
 141 systems (Stips et al., 2016; Bai et al., 2018; Xiao et al., 2020). For two time series  $X_1$  and  
 142  $X_2$ , the rate of information flowing from  $X_2$  to  $X_1$  can be estimated by  $T_{2 \rightarrow 1} = \frac{C_{11}C_{12}C_{2,d1} - C_{12}^2C_{1,d1}}{C_{11}^2C_{22} - C_{11}C_{12}^2}$ ,  
 143 where  $C_{ij}$  is the sample covariance between  $X_i$  and  $X_j$ , and  $C_{i,dj}$  represents the sample  
 144 covariance between  $X_i$  and  $\dot{X}_j$  ( $\dot{X}_j = \frac{dX_j}{dt}$  using Euler forward scheme). When the absolute  
 145 value of  $T_{2 \rightarrow 1}$  is significantly greater than 0,  $X_2$  is the cause of  $X_1$ . Here, the significant test  
 146 of  $T_{2 \rightarrow 1}$  is conducted by the Monte Carlo test approach, in which the 300 random series are  
 147 generated to obtain the 95% confidence interval (CI). If the absolute value of  $T_{2 \rightarrow 1}$  is beyond  
 148 CI, the causal relationship is significant at 95% confidence level (The 95% CI is almost  
 149 unchanged when using more than 300 random series). Moreover, a negative value of  $T_{2 \rightarrow 1}$   
 150 means that  $X_2$  tends to stabilize  $X_1$ , while the positive one implies that  $X_2$  makes  $X_1$  more  
 151 uncertain or more unpredictable.

152 Furthermore, to examine the influence of Kuroshio SSTA on atmospheric circulation, the  
 153 atmospheric general circulation model (AGCM) experiments have been conducted by the

154 Community Atmospheric Model version 6.0 (CAM 6.0; Danabasoglu et al., 2020) with a  
155 horizontal resolution of  $1.9^\circ \times 2.5^\circ$  and 32 vertical levels. The design of numerical  
156 experiment will be demonstrated in Section 3.

### 157 **3. Results**

158 As a persistent medium of air-sea coupled system, SST could be one of the factors  
159 affecting precipitation over Central China in September. To verify the relationship, we  
160 calculate the correlation coefficient between SSTA and precipitation anomaly averaged in  
161 Central China. The results show that the accumulative precipitation here in September has a  
162 significant relationship with SSTA in the Kuroshio region near the ECS (K-ECS:  $125^\circ\text{--}137^\circ\text{E}$ ,  
163  $20^\circ\text{--}34^\circ\text{N}$ ; red box in Fig. 2a) during August. When there is a negative SSTA in K-ECS during  
164 August, the accumulative precipitation over Central China increases in September. The  
165 correlation results between the time series of September precipitation anomaly averaged in  
166 Central China and the August SSTA in K-ECS (Fig. 2a), as well as those between the time  
167 series of August SSTA averaged in K-ECS and the September precipitation anomaly over  
168 Central China (Fig. 2b), both reflect the significant negative correlation relationship. And the  
169 correlation coefficient between these two standardized time series is  $-0.47$  exceeding 95%  
170 confidence level according to the two-tailed Student's  $t$  test (Fig. 2e). Furthermore, to  
171 examine the causality between above two variables, the flow of information from August  
172 SSTA to September precipitation anomaly has been shown in Figs. 2c and 2d. The significant  
173 information flow rate indicates that SSTA in K-ECS during August is a causal factor resulting  
174 in the precipitation anomaly over Central China in September.

175 To explore the anomalous variations of September precipitation over Central China, Fig.  
176 3a displays the anomalous water vapor flux and its divergence integrated from 1000 hPa to  
177 200 hPa in September regressed onto the standardized time series of accumulative  
178 precipitation anomaly averaged over Central China during the same period. Apparently, the  
179 increased precipitation over Central China is accompanied by an anomalous anticyclonic  
180 circulation over the southeastern of China and the Kuroshio region near the East China Sea  
181 during September. Under the action of this anomalous atmospheric circulation, abundant  
182 moisture is transported to Central China by the southerly, which leads to the convergence of  
183 water vapor and further creates a favorable condition for the generation of precipitation here.  
184 Due to the influence of intensive meridional water vapor flux below the middle troposphere,  
185 the moisture convergence over Central China, predominantly composed of meridional  
186 components, is mainly concentrated below 600 hPa (Figs. 3b and 3c).

187 As one of the factors causing the anomalous precipitation over Central China in  
188 September, August SSTA in K-ECS may establish a connection with precipitation by altering  
189 atmospheric circulation. To clarify the specific mechanism of such a time lag, figures 4 and 5  
190 show the variations of circulation and surface heat flux regressed onto the standardized time  
191 series of SSTA during August in K-ECS. Hereafter, the time series of SSTA is multiplied by  
192 a minus sign to visually represent the relationship between it and atmospheric variables. In  
193 August, there is an anomalous cyclonic circulation over K-ECS (Fig. 4a), corresponding to  
194 the active local convection. Thus, the cloud cover increases over K-ECS (Fig. 4a), which  
195 reduces the downward solar radiation here (positive downwards; Fig. 4c). This anomalous

196 atmospheric circulation leads to negative SSTA (Fig. 5a). Concurrent with the negative SSTA  
197 in K-ECS, a positive surface turbulent heat flux anomaly (sum of surface latent and sensible  
198 heat flux) is observed here (Fig. 4e), which indicates a cooling effect of ocean on the overlying  
199 atmosphere. And due to the temporal persistence of SST (Figs. 5a and 5b), this cooling effect  
200 lasts until September (Fig. 4f), which weakens the local convection over K-ECS.  
201 Correspondingly, there is an anomalous descending motion above it in September (Fig. 5c),  
202 which corresponds to the anomalous divergence and the formation of anticyclonic circulation  
203 anomaly at the low troposphere (Fig. 5d). Thus, in September, the cloud cover decreases  
204 over K-ECS (Fig. 4b), and the downward solar radiation increases here (Fig. 4d). Due to the  
205 cooling effect of ocean and its absorption of solar radiation, although the negative SSTA over  
206 K-ECS can persist until September, its amplitude is weakened (Fig. 5b), which also reflects  
207 the feedback of SSTA to the anomalous anticyclonic circulation. Hence, compared with the  
208 situations in September (not shown), the statistical correlation and causal relationships  
209 between August SSTA in K-ECS and precipitation anomaly over Central China are more  
210 significant.

211 Compared with the geopotential height anomaly related to September precipitation over  
212 Central China (Fig. 6a), the anomalous geopotential height affected by August SSTA in K-  
213 ECS is primarily below 600 hPa over K-ECS and the southeast of China (Fig. 6b). To further  
214 confirm the influence of August SSTA in K-ECS on atmospheric circulation in September, the  
215 AGCM experiments has been designed. The global climatological SST with an annual cycle  
216 has been chosen as the forcing condition in control run. The negative SST perturbations in

217 K-ECS have been superimposed onto the SST field of control runs. The resulting combined  
218 SST is then used to force the sensitivity run. Each experiment has been integrated for 30  
219 years. The differences of the last 20 years between the sensitivity and control runs have been  
220 analyzed to verify the atmospheric response to August SSTA in K-ECS. As shown in Figs.  
221 6c and 6d, there is an anticyclonic circulation over K-ECS and the southeast of China, which  
222 is similar to that obtained from reanalysis data (Fig. 6b) except that its coverage area is mainly  
223 concentrated in the west of 125°E. Due to the anomalous anticyclonic circulation in numerical  
224 experiments being biased westward (Figs. 6b, 6c and 6d), there is a westward bias in the  
225 position of anomalous northward transport of water vapor (Figs. 6d and 6e). Hence, the region  
226 of anomalous water vapor convergence in AGCM experiments (blue line in Fig. 6f) differs  
227 from that in reanalysis results (red line in Fig. 6f). In model, the convergence of water vapor  
228 is mainly located to the west of Central China (Fig. 6d and blue line in Fig. 6f), which  
229 corresponds to the area with increased precipitation (not shown). Although there are some  
230 deviations between reanalysis and model results in the spatial patterns of water vapor and  
231 precipitation fields over Central China, the AGCM experiments can capture the features of  
232 atmospheric circulation as shown in reanalysis results, which proves the interpretability of  
233 August SSTA in K-ECS for the anomalous atmospheric circulation transporting moist air to  
234 Central China during September.

235 By comparison, the anomalous atmospheric circulation related to September  
236 precipitation over Central China extends from the lowest pressure level (1000 hPa) to the  
237 upper troposphere (200 hPa), which exhibits the quasi-barotropic structure (Fig. 6a).

238 However, only the portion below 600hPa is affected by August SSTA in K-ECS (Figs. 6b and  
239 6c). Considering that the meridional water vapor transport connected to precipitation is mainly  
240 below 600hPa (Fig. 3c), the reasons for the circulation changes in the upper troposphere  
241 have not been discussed here.

#### 242 **4. Summary and discussions**

243 In this study, we have found that there is a significant negative correlation between  
244 August SSTA in K-ECS and the anomalous precipitation over Central China during  
245 September. And the causality analysis based on Liang–Kleeman information flow theory  
246 (Liang 2014, 2015) has indicated that the former is a cause leading to the September  
247 precipitation anomaly over Central China. The variations of atmospheric circulation related to  
248 both variables have been analyzed to understand the influential mechanism. Due to the  
249 temporal persistence of SST, the cooling effect of negative SSTA in K-ECS on the overlying  
250 atmosphere lasts from August to September, which weakens local convection and  
251 subsequently alters the wind divergence. And the anticyclonic circulation anomaly is formed  
252 above it. As the western part of the anomalous anticyclonic circulation, the southerly carries  
253 out abundant water vapor to Central China, which provides the favorable condition for  
254 generating precipitation here. The anomalous circulation caused by SSTA in K-ECS is chiefly  
255 concentrated below 600 hPa over K-ECS and the southeast of China, which is demonstrated  
256 by the results of AGCM experiments. The anomalous thermal forcing in K-ECS could  
257 establish a close causal relationship with precipitation over Central China during September  
258 by affecting atmospheric circulation, which implies that the August SSTA in K-ECS can be

259 potentially effective as a predictability source for September precipitation over Central China.

260 The oceanic temperature front in the subtropical North Pacific Ocean usually impacts the  
261 variations of westerly and storm track (Chen et al., 2019; Hu et al., 2022, 2023) by changing  
262 the atmospheric baroclinicity. Considering that there is also a strong oceanic temperature  
263 front in K-ECS, in future studies, it is necessary to explore the roles played by Kuroshio SST  
264 front in ECS in altering the atmospheric circulation and its impact on precipitation over China.

265 Besides, as a crucial anticyclonic circulation system in the mid-lower troposphere over  
266 the WNP during warm season, WNPSH is closely related to the precipitation over the eastern  
267 China. As the upstream of WNPSH, the mid-latitude Silk Road wave train is proved to affect  
268 the position and intensity of WNPSH (Chen and Huang, 2012; Guan et al., 2019; Wang et  
269 al., 2023). In general, the influences of the zonal wave trains propagate along the Asian  
270 subtropical westerly jet. The behavior of Silk Road pattern could be modulated by the  
271 meridional position of Asian subtropical westerly jet (Hong and Lu, 2016; Li and Lu, 2017;  
272 Hong et al., 2018), which is usually associated with the thermal and orographic forcing of  
273 Tibetan Plateau (Schiemann et al., 2009; Molnar et al., 2010; Park et al., 2012; Kong and  
274 Chiang, 2020). Taken together, these studies suggest that the Tibetan Plateau forcings might  
275 influence the variations of WNPSH through altering Asian subtropical westerly jet that  
276 modulates the Silk Road pattern. In addition, the heat source over Tibetan Plateau favors the  
277 eastward extension of South Asia high (Wan et al., 2017; Lu et al., 2018; Ge et al., 2019),  
278 which promotes the westward movement of intensified WNPSH by changing the local  
279 atmospheric vorticity (Chen and Zhai, 2016; Guan et al., 2018; Ge et al., 2019; Wei et al.,

280 2019). Based on the above studies, it is indicated the important roles of Tibetan Plateau  
281 forcings in the precipitation over the eastern China. In this study, Central China is located in  
282 the middle and lower reaches of the Yangtze River, which is a part of the eastern China.  
283 Hence, in future studies, it would be necessary to explore whether the precipitation here is  
284 affected by the forcings of Tibetan Plateau. If it is indeed influenced, it will be worthwhile to  
285 analyze the relative importance of Tibetan Plateau forcings, SSTA in K-ECS and Kuroshio  
286 SST front in ECS.

287 In this study, there are some deviations between reanalysis and model results in the  
288 spatial patterns of water vapor and precipitation fields. To improve these differences, the  
289 regional climate model (RCM) is perhaps a good option for the refined simulations. However,  
290 the response of large-scale atmospheric circulation might be limited by the scope of RCM.  
291 Using the atmospheric circulation fields obtained from global models to drive the refined RCM  
292 may be a viable approach. Considering the periodicity and difficulty of conducting related  
293 experiments by using RCM, we will attempt to design AGCM experiments based on refined  
294 RCM for the further explorations.

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### Data Availability Statement

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302 Precipitation provided by NMIC of CMA is downloaded from <http://idata.cma/cmadaas/>.

303 Atmospheric variables from ERA5 are openly available at

304 <https://www.ecmwf.int/en/forecasts/datasets/reanalysis-datasets/era5>. OISST data is

305 derived from [https://www.ncdc.noaa.gov/oisst/optimum-interpolation-sea-surface-](https://www.ncdc.noaa.gov/oisst/optimum-interpolation-sea-surface-temperature-oisst-v21)

306 [temperature-oisst-v21](https://www.ncdc.noaa.gov/oisst/optimum-interpolation-sea-surface-temperature-oisst-v21).

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### Acknowledgments

309 The authors are grateful to the Editor and anonymous reviewers, whose comments and

310 suggestions have greatly improved the manuscript, and are thankful to the public platforms

311 for providing the data listed in section 2. This study was supported by the National Natural

312 Science Foundation of China (U2242206 and 42106036).

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### References

315 Bai, C., R. Zhang, S. Bao, X. S. Liang, W. Guo, 2018: Forecasting the tropical cyclone

316 genesis over the Northwest Pacific through identifying the causal factors in cyclone–climate

317 interactions. *J. Atmospheric Ocean. Technol.*, **35**, 247–259.

318 Bai, H., H. Hu, X. Ren, X.-Q. Yang, Y. Zhang, K. Mao, and Y. Zhao, 2023: The impacts of

319 East China Sea Kuroshio front on winter heavy precipitation events in Southern China. *J.*

320 *Geophys. Res. Atmos.*, **128**, e2022JD037341. <https://doi.org/10.1029/2022JD037341>.

- 321 Bai, H., H. Hu, W. Perrie, and N. Zhang, 2020: On the characteristics and climate effects of  
322 HV-WCP events over the Kuroshio SST front during wintertime. *Clim. Dyn.*, **55**, 2123–2148.
- 323 Barriopedro, D., C. M. Gouveia, R. M. Trigo, and L. Wang, 2012: The 2009/10 drought in  
324 China: possible causes and impacts on vegetation. *J. Hydrometeorol.*, **13**, 1251–1267.
- 325 Chen, G. S., and R. H. Huang, 2012: Excitation mechanisms of the teleconnection patterns  
326 affecting the July precipitation in Northwest China. *J. Clim.*, **25**, 7834–7851.
- 327 Chen, Q. Y., H. B. Hu, X. J. Ren, and X.-Q. Yang, 2019: Numerical Simulation of Midlatitude  
328 Upper - Level Zonal Wind Response to the Change of North Pacific Subtropical Front  
329 Strength, *J. Geophys. Res. Atmos.*, **124**, 4891–4912.
- 330 Chen, Y., and P. Zhai, 2016: Mechanisms for concurrent low-latitude circulation anomalies  
331 responsible for persistent extreme precipitation in the Yangtze River Valley. *Clim*  
332 *Dyn.*, **47**, 989–1006.
- 333 Danabasoglu, G., J. F. Lamarque, J. Bacmeister, D. A. Bailey, A. K. DuVivier, J. Edwards,  
334 L. K. Emmons, J. Fasullo, R. Garcia, A. Gettelman, C. Hannay, M. M. Holland, W. G. Large,  
335 P. H. Lauritzen, D. M. Lawrence, J. T. Lenaerts, K. Lindsay, W. H. Lipscomb, M. J. Mills,  
336 R. Neale, K. W. Oleson, B. Otto-Bliesner, A. S. Phillips, W. Sacks, S. Tilmes, L. van  
337 Kampenhout, M. Vertenstein, A. Bertini, J. Dennis, C. Deser, C. Fischer, B. Fox-Kemper,  
338 J. E. Kay, D. Kinnison, P. J. Kushner, V. E. Larson, M. C. Long, S. Mickelson, J. K. Moore,  
339 E. Nienhouse, L. Polvani, P. J. Rasch, and W. G. Strand, 2020: The Community Earth  
340 System Model Version 2 (CESM2). *J. Adv. Model. Earth Sy.*, **12**, 1–35.
- 341 Fang, Y.H., K.-Q. Chen, H.-S. Chen, S.-Q. Xu, X. Geng, T.-Y. Li, F.-D. Teng, X.-Y. Zhou and

- 342 Y.-G. Wang, 2018: The remote responses of early summer cold vortex precipitation in  
343 Northeastern China to the precedent sea surface temperatures. *Atmos. Res.*, **214**, 399–  
344 409.
- 345 Feng, J., and X.-C. Wang, 2018: Impact of two types of La Niña on boreal autumn rainfall  
346 around Southeast Asia and Australia. *Atmos. Ocean. Sci. Lett.*, **11**, 1–6.
- 347 Gan, B., Y.-O. Kwon, T. M. Joyce, K. Chen and L. Wu, 2019: Influence of the Kuroshio  
348 Interannual Variability on the Summertime Precipitation over the East China Sea and  
349 Adjacent Area. *J. Clim.*, **32**, 2185–2205.
- 350 Ge, J., Q. You, and Y. Zhang, 2019: Effect of Tibetan Plateau heating on summer extreme  
351 precipitation in eastern China. *Atmos. Res.*, **218**, 364–371.
- 352 Geng, Y., H.-L. Ren, X. Ma, S. Zhao, and Y. Nie, 2022: Responses of East Asian Climate to  
353 SST Anomalies in the Kuroshio Extension Region during Boreal Autumn. *J. Clim.*, **35**,  
354 7007–7023.
- 355 Gu, W., L. Wang, W. Li, L. Chen, and C. Sun, 2014: Influence of the Tropical Pacific East–  
356 West Thermal Contrast on the Autumn Precipitation in South China. *Int. J. Climatol.*, **35**,  
357 1543–1555.
- 358 Guan, W., X. Ren, W. Shang, and H. Hu, 2018: Subseasonal Zonal Oscillation of the Western  
359 Pacific Subtropical High during Early Summer. *J. Meteorol. Res.*, **32**, 768–780.
- 360 Guan, W. N., H. B. Hu, X. J. Ren, and X.-Q. Yang, 2019: Subseasonal zonal variability of the  
361 western Pacific subtropical high in summer: climate impacts and underlying mechanisms.  
362 *Clim. Dyn.*, **53**, 3325–3344.

- 363 Hersbach, H., B. Bell, P. Berrisford, G. Biavati, A. Horányi, J. Muñoz Sabater, J. Nicolas, C.  
364 Peubey, R. Radu, I. Rozum, D. Schepers, A. Simmons, C. Soci, D. Dee, and J-N. Thépaut,  
365 2019: ERA5 monthly averaged data on pressure levels from 1979 to present. Copernicus  
366 Climate Change Service (C3S) Climate Data Store (CDS). (Accessed on < DD-MMM-  
367 YYYY >), 10.24381/cds.6860a573.
- 368 Hong, C.-C., Y. K. Wu, and T. Li, 2016: Influence of climate regime shift on the interdecadal  
369 change in tropical cyclone activity over the Pacific Basin during the middle to late 1990s.  
370 *Clim. Dyn.*, **47**, 2587– 2600.
- 371 Hong, X., and R. Lu, 2016: The meridional displacement of the summer Asian jet, Silk Road  
372 pattern, and tropical SST anomalies. *J. Climate*, **29**, 3753–3766.
- 373 Hong, X., R. Lu, and S. Li, 2018: Asymmetric relationship between the meridional  
374 displacement of the Asian westerly jet and the Silk Road pattern. *Adv. Atmos. Sci.*, **35**,  
375 389–396.
- 376 Hu, C., C. Zhang, S. Yang, D. Chen, and S. He, 2017: Perspective on the northwestward  
377 shift of autumn tropical cyclogenesis locations over the western North Pacific from shifting  
378 ENSO. *Clim. Dyn.*, **51**, 2455–2465.
- 379 Hu, C., D. Chen, G. Huang, and S. Yang, 2018: Dipole types of autumn precipitation variability  
380 over the subtropical East Asia-western Pacific modulated by shifting ENSO. *Geophys. Res.*  
381 *Lett.*, **45**, 9123–9130.
- 382 Hu, H. B., W. X. Chen, X.-Q. Yang, Y. H. Zhao, H. K. Bai, and K. F. Mao, 2022: The Mode-  
383 Water-Induced Interannual Variation of the North Pacific Subtropical Countercurrent and

- 384 the Corresponding Winter Atmospheric Anomalies, *Geophys. Res. Lett.*, **49**,  
385 e2022GL100968. <http://doi.org/10.1029/2022GL100968>.
- 386 Hu, H. B., Y. Zhao, X.-Q. Yang, S. Jiang, K. Mao, and H. Bai, 2023: The influences of the  
387 multi-scale sea surface temperature anomalies in the North Pacific on the jet stream in  
388 winter. *J. Geophys. Res. Atmos.*, **128**, e2022JD038036.  
389 <https://doi.org/10.1029/2022JD038036>.
- 390 Huang, B., C. Liu, V. Banzon, E. Freeman, G. Graham, B. Hankins, T. Smith, and H.-M.  
391 Zhang, 2021: Improvements of the Daily Optimum Interpolation Sea Surface Temperature  
392 (DOISST) Version 2.1. *J. Clim.*, **34**, 2923-2939.
- 393 Kong, W., and J. C. H. Chiang, 2020: Interaction of the westerlies with the Tibetan Plateau in  
394 determining the Mei-Yu termination. *J. Climate*, **33**, 339–363.
- 395 Li, C., and T. Zhao, 2019: Seasonal Responses of Precipitation in China to El Niño and  
396 Positive Indian Ocean Dipole Modes. *Atmosphere*, **10**, 372.  
397 <https://doi.org/10.3390/atmos10070372>.
- 398 Li, X., and R. Lu, 2017: Extratropical factors affecting the variability in summer precipitation  
399 over the Yangtze River Basin, China. *J. Climate*, **30**, 8357–8374.
- 400 Liang, X., 2014: Unraveling the cause-effect relation between time series. *Phys. Rev. E.*, **90**,  
401 052150. <https://doi.org/10.1103/PhysRevE.90.052150>.
- 402 Liang, X., 2015: Normalizing the causality between time series. *Phys. Rev. E.*, **92**, 022126.  
403 <https://doi.org/10.1103/PhysRevE.92.022126>
- 404 Lin, Z., J. Zhu, W. Hua, and G. Fan, 2021: Impact of the August Asian–Pacific Oscillation on

- 405 Autumn Precipitation in Central Eastern China. *Asia-Pac. J. Atmos. Sci.*, **57**, 181–190.
- 406 Liu, B., C. Zhu, S. Ma, and Y. Yan, 2022: Combined Effects of Tropical Indo-Pacific–Atlantic  
407 SST Anomalies on Record-Breaking Floods over Central-North China in September 2021.  
408 *J. Clim.*, **35**, 6191–6205.
- 409 Liu, B., Y. Yan, C. Zhu, S. Ma, and J. Li, 2020: Record - breaking Meiyu rainfall around the  
410 Yangtze River in 2020 regulated by the subseasonal phase transition of the North Atlantic  
411 Oscillation. *Geophys. Res. Lett.*, **47**, e2020GL090342. [https://doi.](https://doi.org/10.1029/2020GL090342)  
412 [org/10.1029/2020GL090342](https://doi.org/10.1029/2020GL090342).
- 413 Lu, M. M., S. Yang, Z. N. Li, B. He, S. He, and Z. Q. Wang, 2018: Possible effect of the  
414 Tibetan Plateau on the “upstream” climate over West Asia, North Africa, South Europe and  
415 the North Atlantic. *Clim. Dyn.*, **51**, 1485–1498.
- 416 Matsumura, S., and T. Horinouchi, 2023: Decadal Shift in Summer Precipitation Variability  
417 over East Asia in the Mid-2000s and Wave Propagation toward North America. *J. Clim.*,  
418 **36**, 2483–2496.
- 419 Minobe, S., and S. Takebayashi, 2015: Diurnal precipitation and high cloud frequency  
420 variability over the Gulf Stream and over the Kuroshio. *Clim. Dyn.*, **44**, 2079–2095.
- 421 Molnar, P., W. R. Boos, and D. S. Battisti, 2010: Orographic controls on climate and  
422 paleoclimate of Asia: Thermal and mechanical roles for the Tibetan Plateau. *Annu. Rev.*  
423 *Earth Planet. Sci.*, **38**, 77–102.
- 424 Niu, N., and J. P. Li, 2008: Interannual variability of autumn precipitation over South China  
425 and its relation to atmospheric drought and SST anomalies. *Adv. Atmos. Sci.*, **25**, 117–125.

- 426 Park, H.-S., J. C. H. Chiang, and S. Bordoni, 2012: The mechanical impact of the Tibetan  
427 Plateau on the seasonal evolution of the South Asian monsoon. *J. Climate*, **25**, 2394–2407.
- 428 Qi, L., Y. Ji, and W. Zhang, 2021: Indispensable role of the Madden-Julian oscillation in the  
429 2019 extreme autumn drought over eastern China. *J. Geophys. Res. Atmos.*, **126**,  
430 e2020JD034123. <https://doi.org/10.1029/2020JD034123>.
- 431 Qin, M., D. Li, A. Dai, W. Hua, and H. Ma, 2018: The influence of the Pacific decadal  
432 oscillation on north Central China precipitation during boreal autumn. *Int. J.*  
433 *Climatol.*, **38**, 821–831.
- 434 Reynolds, R. W., T. M. Smith, C. Liu, D. B. Chelton, K. S. Casey, and M. G. Schlax, 2007:  
435 Daily high-resolution-blended analyses for sea surface temperature. *J. Clim.*, **20**, 5473–  
436 5496.
- 437 Sasaki, Y. N., S. Minobe, T. Asai, and M. Inatsu, 2012: Influence of the Kuroshio in the East  
438 China Sea on the Early Summer (Baiu) Rain. *J. Clim.*, **25**, 6627–6645.
- 439 Sasaki, Y.N., and Y. Yamada, 2018: Atmospheric response to interannual variability of sea  
440 surface temperature front in the East China Sea in early summer. *Clim. Dyn.*, **51**, 2509–  
441 2522.
- 442 Schiemann, R., D. Lüthi, and C. Schär, 2009: Seasonality and interannual variability of the  
443 westerly jet in the Tibetan Plateau region. *J. Climate*, **22**, 2940–2957.
- 444 Shen, H., S. He, and H. Wang, 2019: Effect of summer Arctic sea ice on the reverse August  
445 precipitation anomaly in Eastern China between 1998 and 2016. *J. Clim.*, **32**, 3389–3407.

- 446 Stips, A., D. Macias, C. Coughlan, E. Garciagoriz, X. Liang, 2016: On the causal structure  
447 between CO<sub>2</sub> and global temperature. *Sci. Rep.*, **6**, 21691.  
448 <https://doi.org/10.1038/srep21691>.
- 449 Wan, B., Z. Gao, F. Chen, and C. Lu, 2017: Impact of Tibetan Plateau surface heating on  
450 persistent extreme precipitation events in southeastern China. *Mon. Wea. Rev.*, **145**,  
451 3485–3505.
- 452 Wang, M., Q. Gu, X. Jia, and J. Ge, 2019: An assessment of the impact of PDO on autumn  
453 droughts in North China based on the PDSI. *Int. J. Climatol.*, **39**, 5338–5350.
- 454 Wang, Y. N., H. B. Hu, X. J. Ren, X.-Q. Yang, and K. F. Mao, 2023: Significant Northward  
455 Jump of the Western Pacific Subtropical High: The Interannual Variability and Mechanisms.  
456 *J. Geophys. Res. Atmos.*, **128**, e2022JD037742. <http://doi.org/10.1029/2022JD037742>.
- 457 Wei, W., R. Zhang, M. Wen, S. Yang, and W. Li, 2019: Dynamic effect of the South Asian  
458 high on the interannual zonal extension of the western North Pacific subtropical high. *Int.*  
459 *J. Climatol.*, **39**, 5367–5379.
- 460 Wu, L., B. Wang, and S. Geng, 2005: Growing typhoon influence on East Asia. *Geophys.*  
461 *Res. Lett.*, **32**, L18703, <https://doi.org/10.1029/2005GL022937>.
- 462 Xiao, M., Z., Q. Zhang, and V. P. Singh, 2015: Influences of ENSO, NAO, IOD and PDO on  
463 seasonal precipitation regimes in the Yangtze River basin, China. *Int. J. Climatol.*, **35**,  
464 3556–3567.

- 465 Xiao H., F. Zhang, L. Miao, X. S. Liang, K. Wu, R. Liu, 2020: Long-term trends in Arctic  
466 surface temperature and potential causality over the last 100 years. *Clim. Dyn.*, **55**, 1443–  
467 1456.
- 468 Xu H.-L., J. Feng, and C. Sun, 2013: Impact of Preceding Summer North Atlantic Oscillation  
469 on Early Autumn Precipitation over Central China. *Atmos. Ocean. Sci. Lett.*, **6**, 417–422.
- 470 Xu, K., C. Zhu, and W. Wang, 2016: The Cooperative Impacts of the El Niño-Southern  
471 Oscillation and the Indian Ocean Dipole on the Interannual Variability of Autumn Rainfall  
472 in China. *Int. J. Climatol.*, **36**, 1987–1999.
- 473 Xu, M., H. Xu, and H. Ren, 2018: Influence of Kuroshio SST front in the East China Sea on  
474 the climatological evolution of Meiyu rainband. *Clim. Dyn.*, **50**, 1243–1266.
- 475 Yang, J., D. Gong, W. Wang, M. Hu, and R. Mao, 2012: Extreme drought event of 2009/2010  
476 over southwestern China. *Meteorol. Atmos. Phys.*, **115**, 173–184.
- 477 Yeh, T.-C., S.-Y. Tao, and M.-T. Li, 1958: The abrupt change of circulation over the Northern  
478 Hemisphere during June and October (in Chinese). *Acta Meteor. Sin.*, **29**, 249–263.
- 479 Ying, K., X. Zheng, T. Zhao, C. S. Frederiksen, and X.-W. Quan, 2017: Identifying the  
480 predictable and unpredictable patterns of spring-to-autumn precipitation over eastern  
481 China. *Clim. Dyn.*, **48**, 3183–3206.
- 482 Yuan, C., and D. Wang, 2019: Interdecadal Variations in El Niño-Southern Oscillation  
483 Impacts on the Autumn Precipitation in the Eastern China. *Int. J. Climatol.*, **39**, 5316–5326.
- 484 Zhang, W., F.-F. Jin, J. X. Zhao, L. Qi, and H.-L. Ren, 2013: The possible influence of a non-  
485 conventional El Niño on the severe autumn drought of 2009 in Southwest China. *J. Clim.*,

486 **26**, 8392–8405.

487 Zhang, W., F.-F. Jin, and A. Turner, 2014: Increasing autumn drought over southern China  
488 associated with ENSO regime shift. *Geophys. Res. Lett.*, **41**, 4020–4026.

489 Zhang, Q., P. Sun, V. P. Singh, and X. Chen, 2012: Spatial-temporal precipitation changes  
490 (1956–2000) and their implications for agriculture in China. *Global Planet. Change*, **82–83**,  
491 86–95.

492 Zhang, Q., Y. Yao, Y. Wang, S. Wang, J. Wang, Yang, J., J. Wang, Y. Li, J. Shang, and W.  
493 Li, 2019: Characteristics of drought in Southern China under climatic warming, the risk,  
494 and countermeasures for prevention and control. *Theor. Appl. Climatol.*, **136**, 1157–1173.

495 Zhou, L., S. Wang, Y. Chi, and J. Wang, 2018: Drought Impacts on Vegetation Indices and  
496 Productivity of Terrestrial Ecosystems in Southwestern China During 2001-2012. *Chinese*  
497 *Geogr. Sci.*, **28**, 784–796.

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510 box in (b) represents the Central China (CC) domain. (e) Standard deviation of precipitation  
511 averaged in CC during autumn and each month in this season.

512 **Fig. 2** (a) Correlation coefficient between the distribution of August SSTA and the time series  
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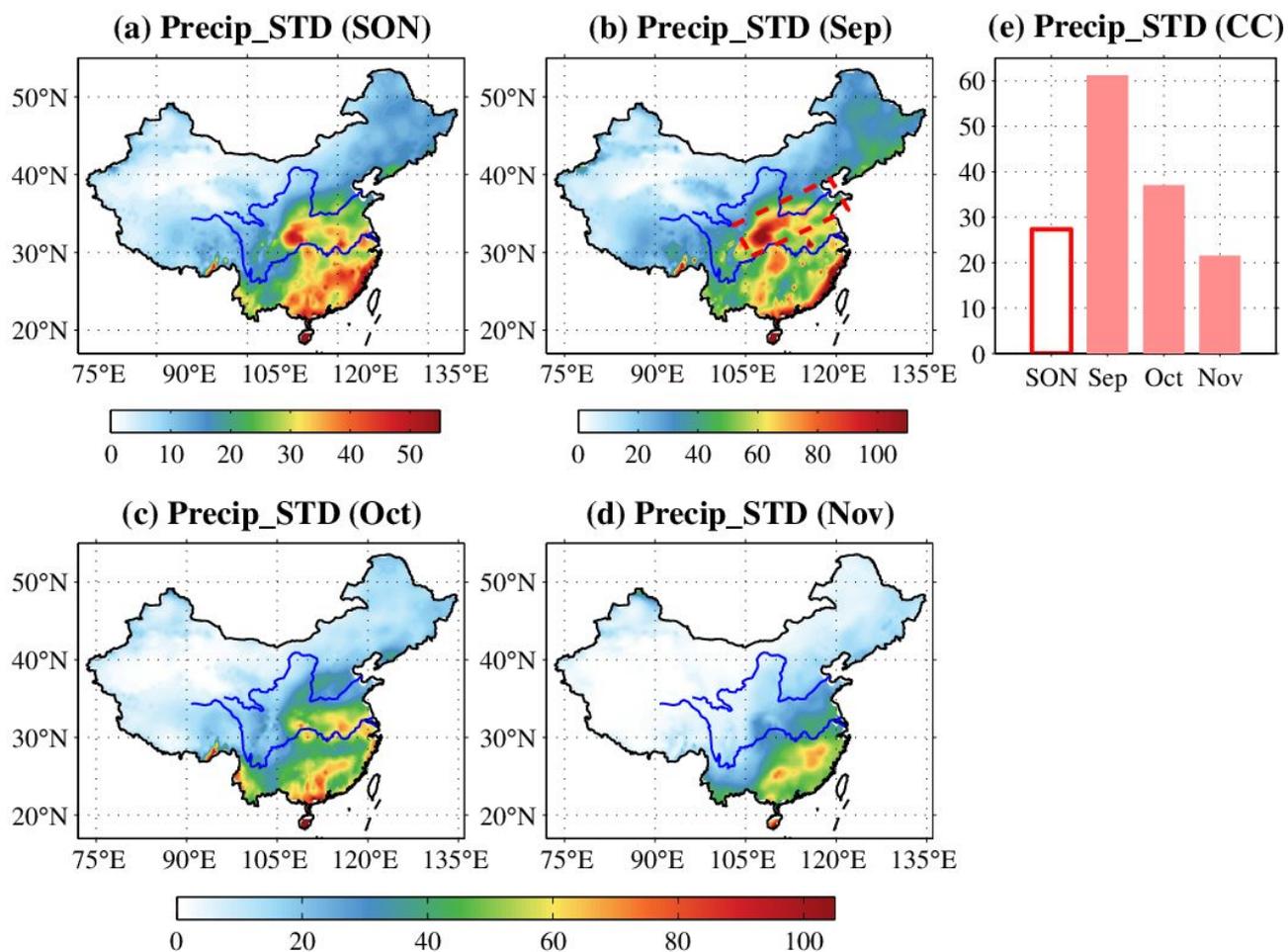
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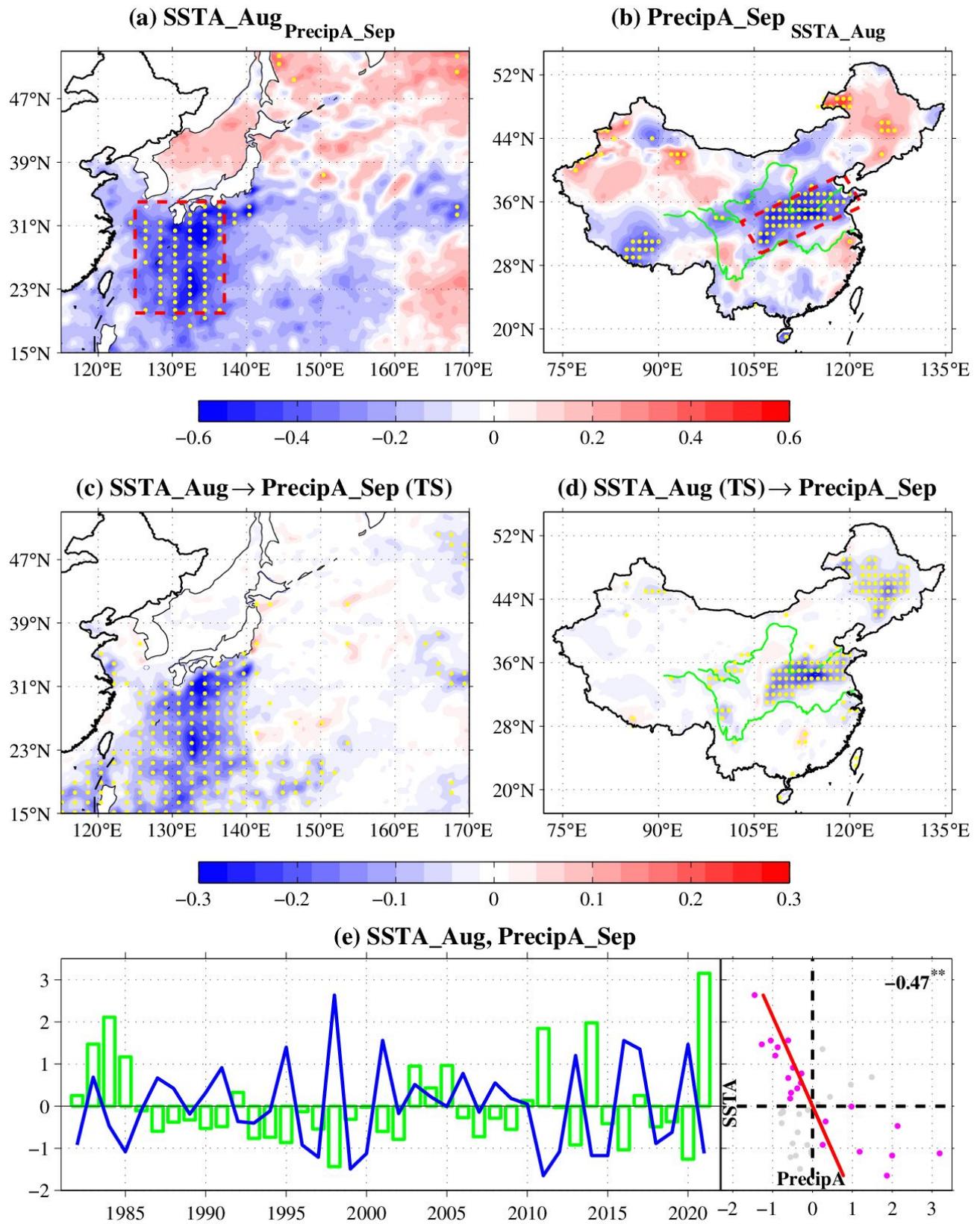
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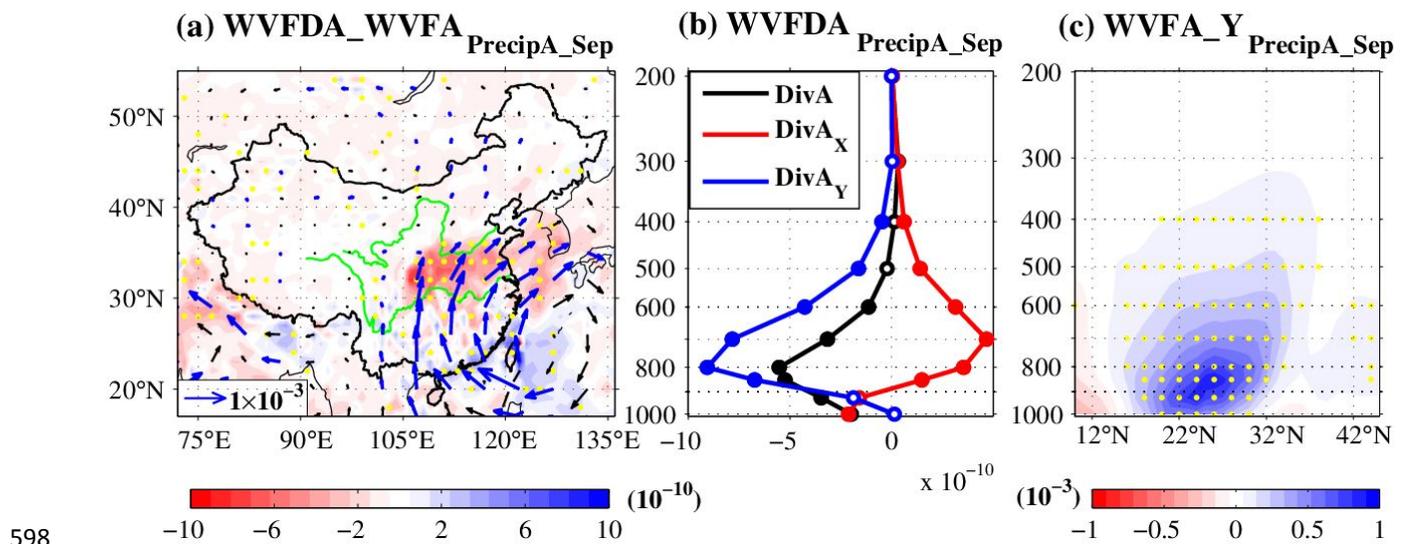


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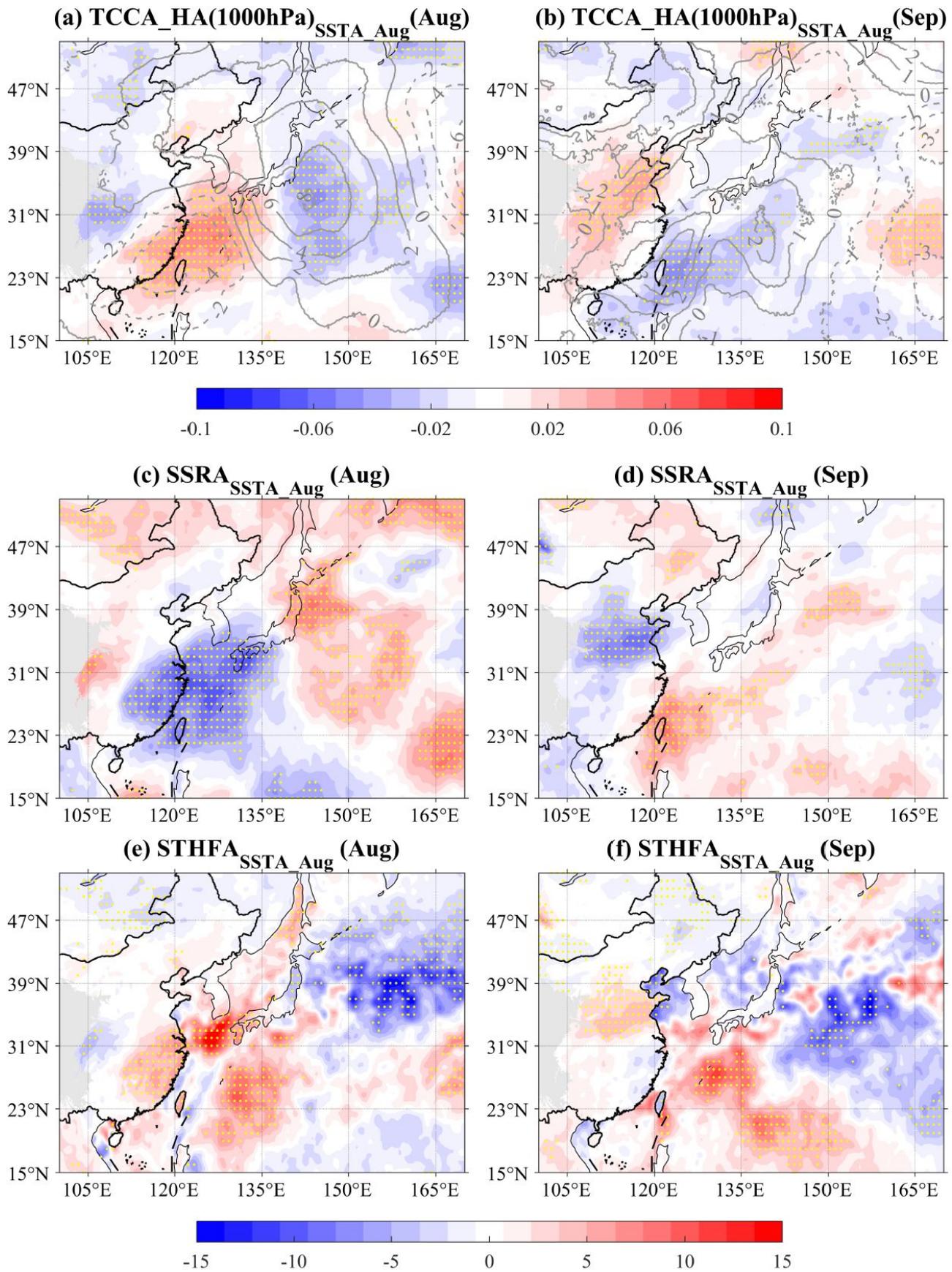
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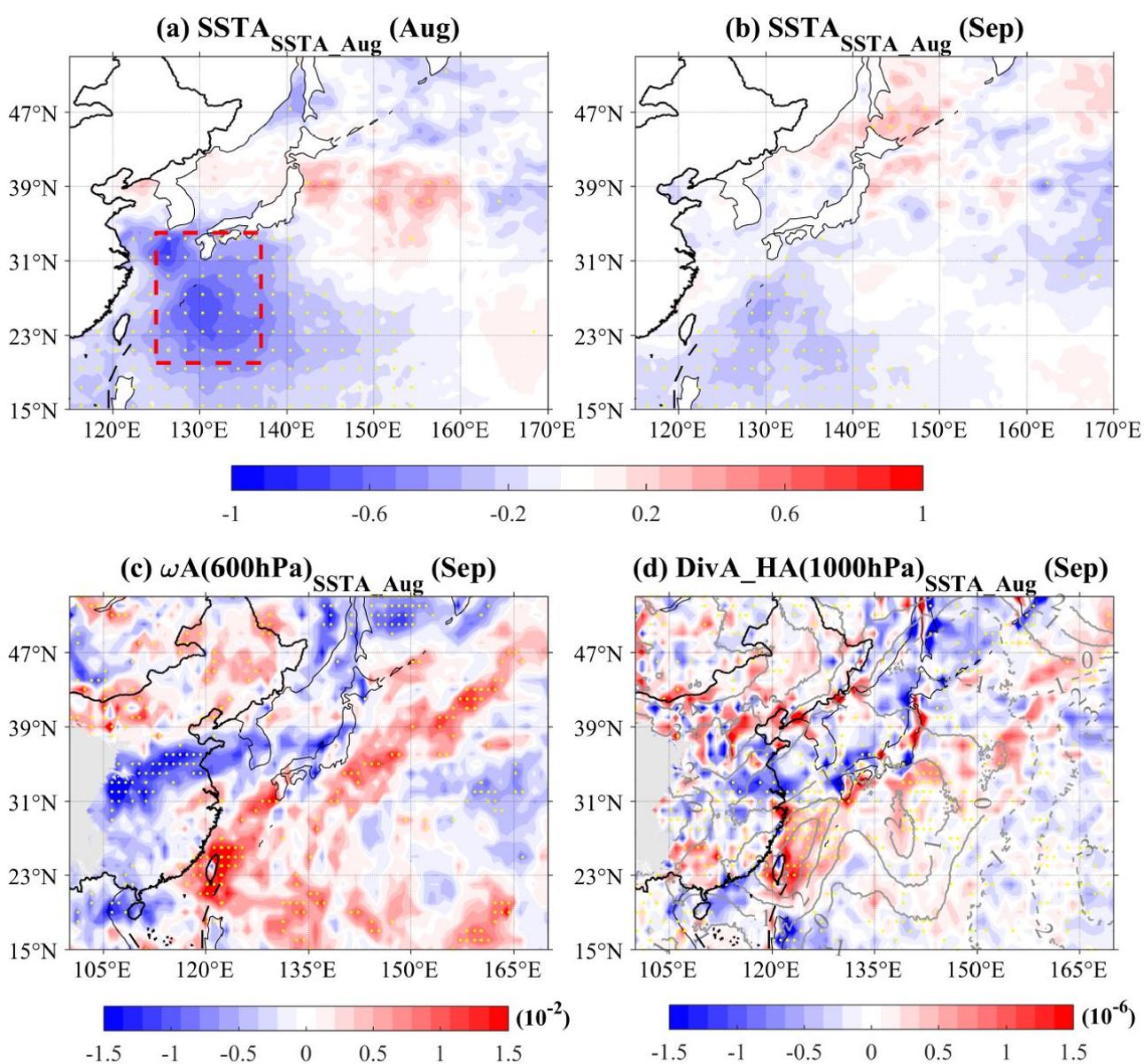
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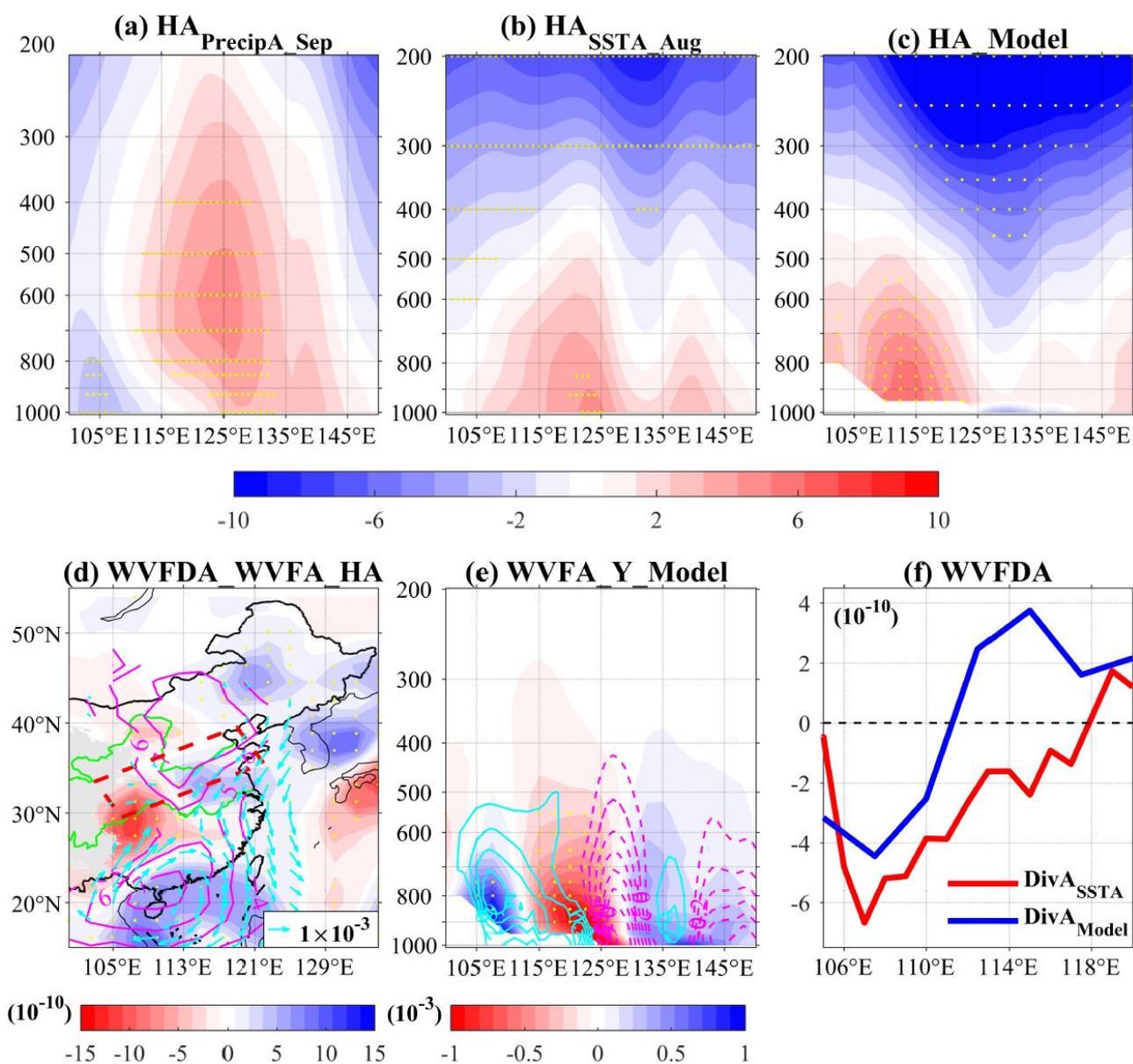
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